

Uncertainty Quantification for Geological Carbon Sequestration

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Workshop on "Opportunities and Challenges in Uncertainty
Quantification for Complex Interacting systems"
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Carbon Management

Carbon Intensity

Energy Intensity

$$\text{Net } C = \underbrace{(C/E) * (E/GDP)}_{= C/GDP} * GDP - \text{Sinks}$$

Carbon sequestration: Capture

Carbon Intensity of Economy

Natural sinks, and artificial sinks (carbon sequestration)

Carbon Sequestration is a MUST

•Fossil fuels will remain the primary energy source well into this century

•Natural sinks are going down because of deforestation and change in agriculture patterns

Desired

$$Net\ C = (C/E) * (E/GDP) * GDP - Sinks$$

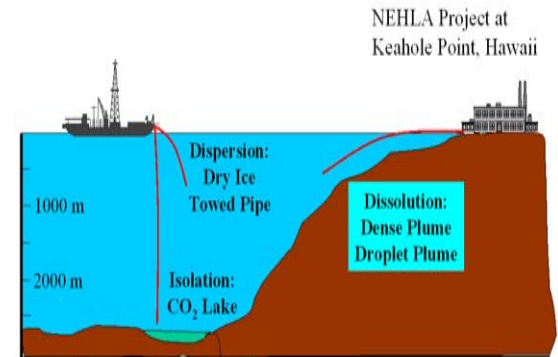
Present

•Energy efficiency improvements are not enough to offset the increased energy demand

•As developing countries' economies expand, worldwide energy consumption will continue its rapid growth

Carbon Sequestration Approaches

- Ocean (>1400 Gt Est. Capacity)
 - Direct injection
 - Enhancement of natural processes
 - Largest capacity, but least understood
- Terrestrial (>10 Gt Est. Capacity)
 - Crops & land management
 - Soil improvement
 - C-species selection
- Geologic (>1000 Gt Est. Capacity)
 - Oil & gas reservoirs
 - Deep saline aquifers
 - Unmineable coal beds



Geological Sequestration Presents a Viable Option

- Past experience
 - Recent: Natural gas storage, groundwater remediation
 - Century-long: Oil/gas production, groundwater resources management
 - Geological time: Natural CO₂ reservoirs (e.g., Bravo Dome, NM)
- Economic benefits that offset the cost of sequestration
 - Enhanced oil recovery from oil reservoirs
 - Increased gas production from natural gas reservoirs
 - Enhanced methane production from coal beds
- Worldwide availability and in close proximity of geological media to CO₂ sources

Basic Research Questions

- The science of geological sequestration is just starting

- Many R &D questions remain:

Capacity

Injectivity

Impact to reservoir

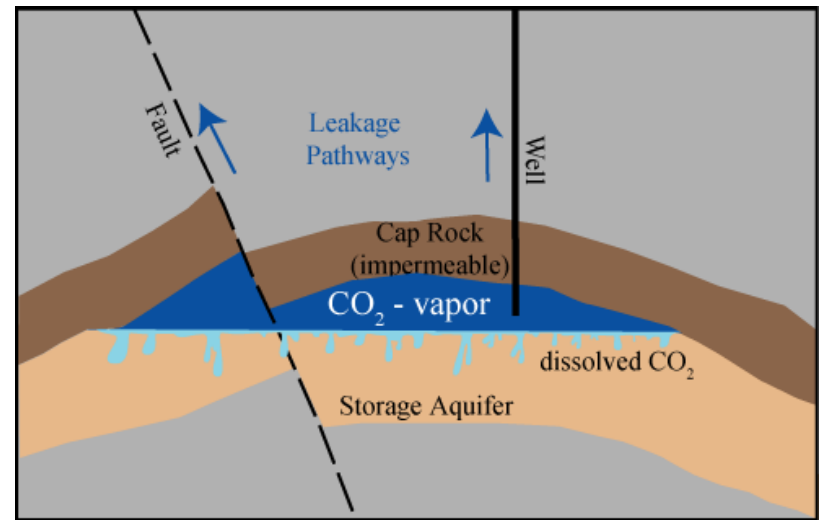
Caprock integrity

Leakage pathway and rate

Monitoring

Performance assessment

Risk assessment



EOR/EGR vs. Sequestration

- **Although industrial EOR experiences exist, there are major differences between EOR and Sequestration:**
 - The current EOR minimizes the amount of CO₂ used while the goal of seq. is to store as much CO₂ as possible
 - EOR is a short term process (of several years) while seq. is at the scale of 100s to 1000s years
 - Long term performance assessment required for sequestration --- EOR has an industrial experience of 40 years (still a short timeframe for seq.)
 - Higher-confidence predictive and monitoring tools are needed for sequestration
- **New EOR strategies are needed if sequestration is the goal**

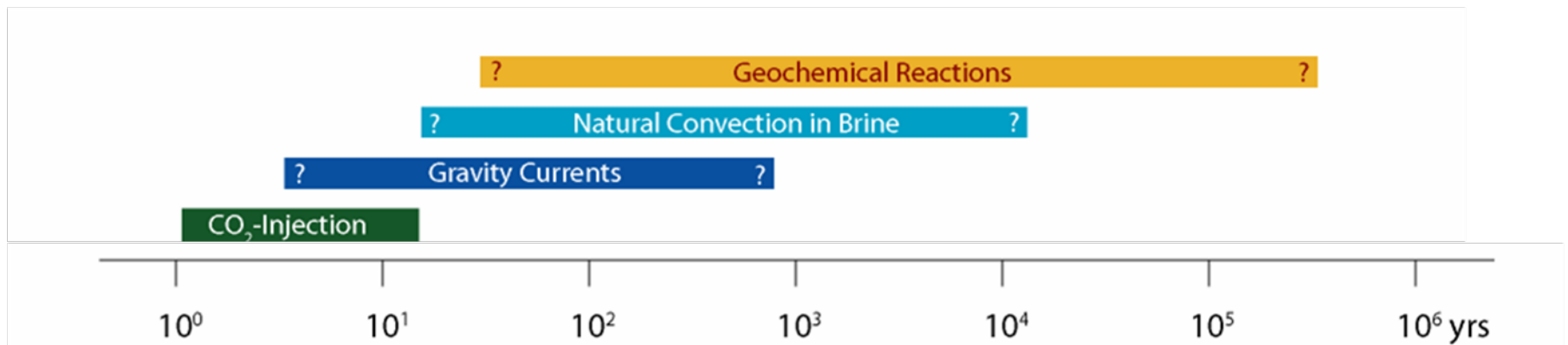
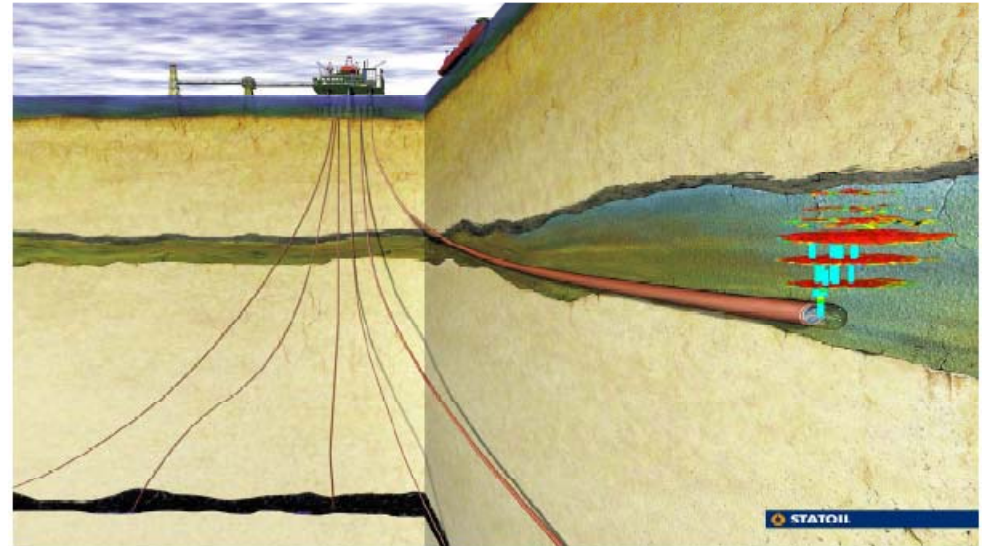
Saline Aquifers

Features:

- Viscous fingering
- Gravity segregation
- Capillary entrapment
- Convective mixing
- Geochemistry

Time scale estimates

Sleipner CO2 Injection



Complexities and Uncertainties

– A multiplicity of length scales

- From atomistic and microscopic, to macroscopic and to field-scale

– Large timescale range of interest

- From picoseconds to millennia

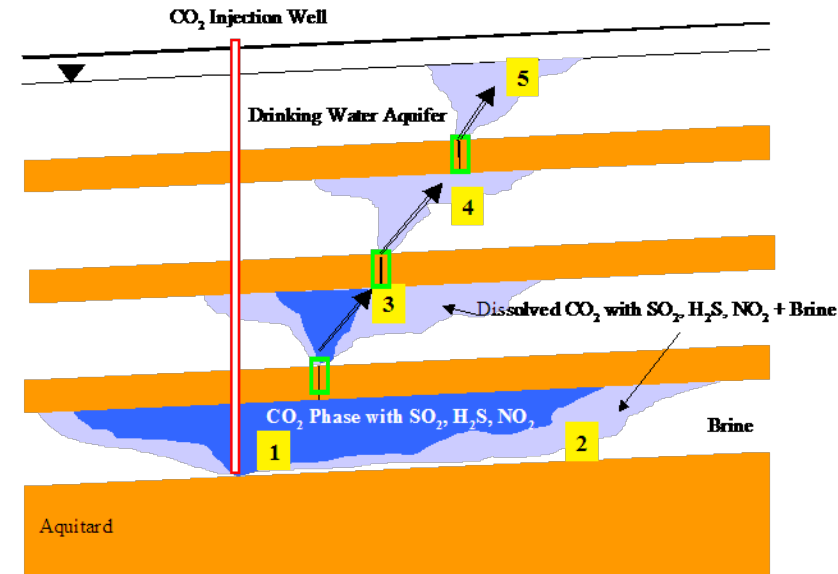
– Coupled processes

- Fluid flow, geomechanics, geochemistry, and heat transfer

– Various components

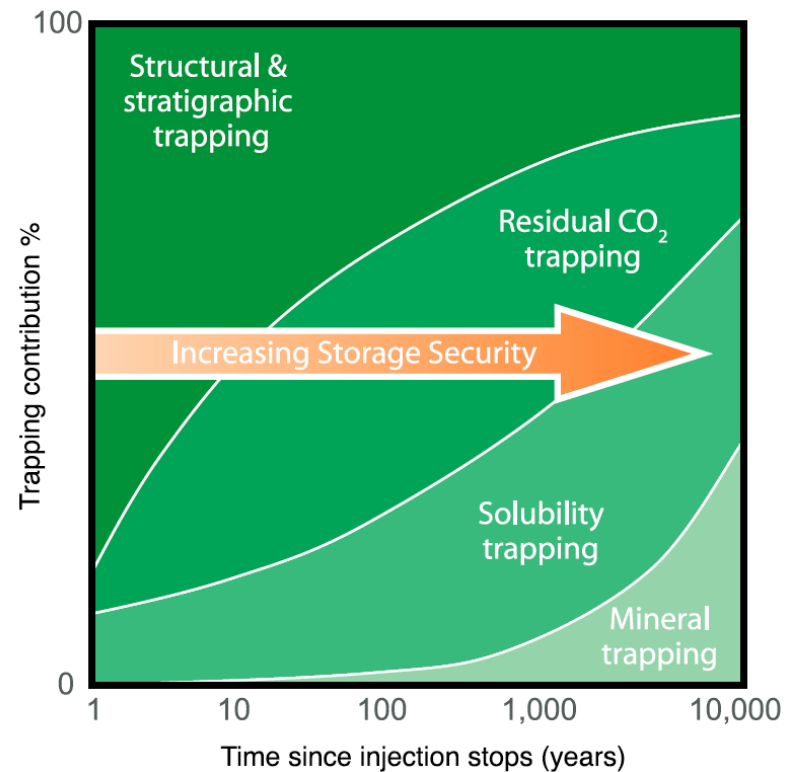
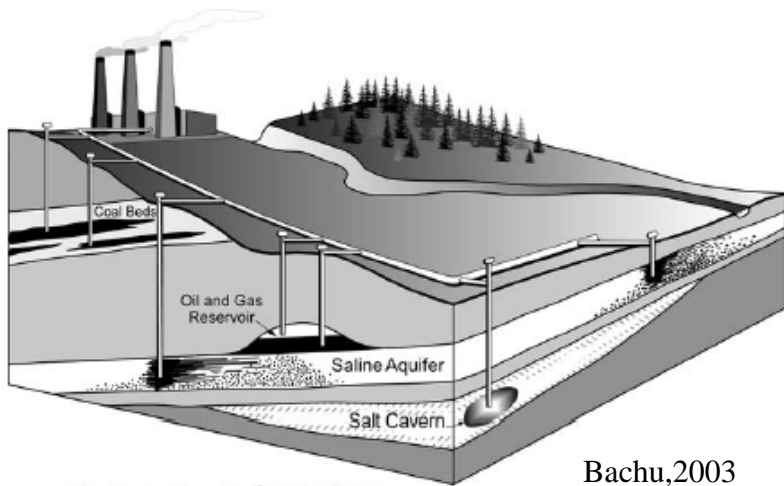
- Reservoir/aquifer, caprock, overburdens, faults, and wells

– Spatial variabilities and poor knowledge of them



CO₂ Trapping Mechanisms

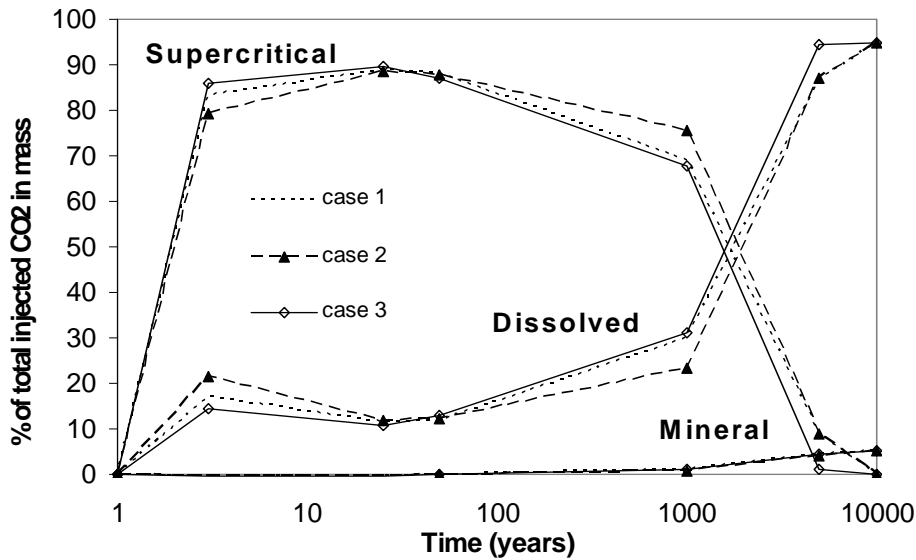
- Structural trapping
- Residual gas trapping
- Solubility trapping
- Mineral trapping



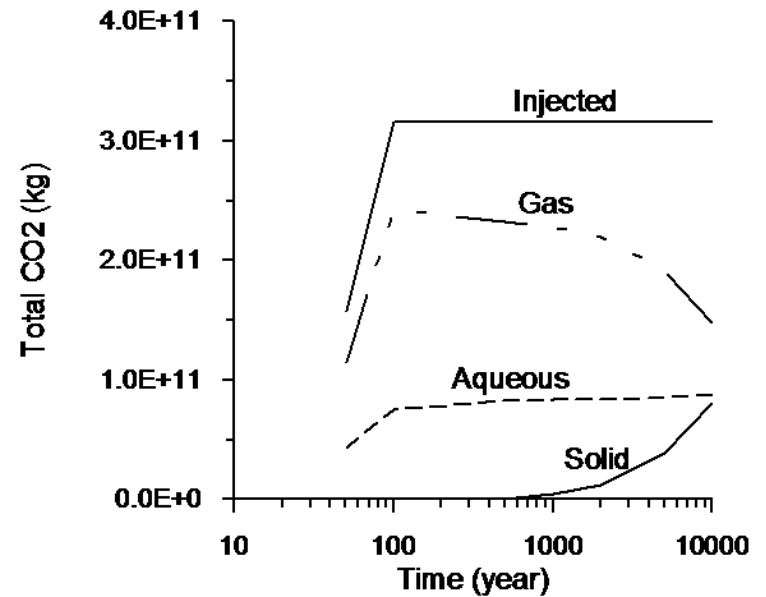
(2005 IPCC Special Report on Carbon Dioxide Capture and Storage;
<http://www.ipcc.ch/activity/srccs/index.htm>)

Fate of Injected CO₂

Sleipner (Audigane et al, 2007)



Gulf Coast Sandstone (Xu et al., 2003)



Hydrologic Properties

Aquifer thickness	100 m
Permeability	10^{-13} m^2
Porosity	0.30
Compressibility	$4.5 \times 10^{-10} \text{ Pa}^{-1}$
Temperature	$75 \text{ }^\circ\text{C}$
Pressure	200 bar
Salinity	0.06 (mass fraction)
CO ₂ injection rate	90 kg/s
Relative permeability	
Liquid (van Genuchten, 1980):	
$k_{rl} = \sqrt{S^*} \left\{ 1 - \left(1 - [S^*]^{1/m} \right)^m \right\}^2$	$S^* = (S_1 - S_{lr}) / (1 - S_{lr})$
irreducible water saturation exponent	$S_{lr} = 0.30$ $m = 0.457$
Gas (Corey, 1954):	
$k_{rg} = (1 - \hat{S})^2 (1 - \hat{S}^2)$	$\hat{S} = \frac{(S_1 - S_{lr})}{(S_1 - S_{lr} - S_{gr})}$
irreducible gas saturation	$S_{gr} = 0.05$
Capillary pressure	
van Genuchten (1980)	
$P_{cap} = -P_0 \left([S^*]^{1/m} - 1 \right)^{1-m}$	$S^* = (S_1 - S_{lr}) / (1 - S_{lr})$
irreducible water saturation exponent	$S_{lr} = 0.00$ $m = 0.457$
strength coefficient	$P_0 = 19.61 \text{ kPa}$

Similar to Gulf Coast Sandstone (Xu et al., 2003)

Geochemical Properties

Mineral	Vol. % Of solid	A (cm ² /g)	Parameters for kinetic rate law								
			Neutral mechanism		Acid mechanism			Base mechanism			
			k ₂₅ (mol/m ² /s)	E _a (KJ /mol)	k ₂₅	E _a	n(H ⁺)	k ₂₅	E _a	n(H ⁺)	
<i>Primary:</i>											
Quartz	57.888	9.8	1.023×10 ⁻¹⁴	87.7							
Kaolinite	2.015	151.6	6.918×10 ⁻¹⁴	22.2	4.898×10 ⁻¹²	65.9	0.777	8.913×10 ⁻¹⁸	17.9	-0.472	
Calcite	1.929		Assumed at equilibrium								
Illite	0.954	151.6	1.660×10 ⁻¹³	35	1.047×10 ⁻¹¹	23.6	0.34	3.020×10 ⁻¹⁷	58.9	-0.4	
Oligoclase	19.795	9.8	1.445×10 ⁻¹³	69.8	2.138×10 ⁻¹¹	65	0.457				
K-feldspar	8.179	9.8	3.890×10 ⁻¹³	38	8.710×10 ⁻¹¹	51.7	0.5	6.310×10 ⁻¹²	94.1	-0.823	
Na-smectite	3.897	151.6	1.660×10 ⁻¹³	35	1.047×10 ⁻¹¹	23.6	0.34	3.020×10 ⁻¹⁷	58.9	-0.4	
Chlorite	4.556	9.8	3.02×10 ⁻¹³	88	7.762×10 ⁻¹²	88	0.5				
Hematite	0.497	12.9	2.512×10 ⁻¹⁵	66.2	4.074×10 ⁻¹⁰	66.2	1				
<i>Secondary:</i>											
Magnesite		9.8	4.571×10 ⁻¹⁰	23.5	4.169×10 ⁻⁷	14.4	1				
Dolomite		9.8	2.951×10 ⁻⁸	52.2	6.457×10 ⁻⁴	36.1	0.5				
Low-albite		9.8	2.754×10 ⁻¹³	69.8	6.918×10 ⁻¹¹	65	0.457	2.512×10 ⁻¹⁶	71	-0.572	
Siderite		9.8	1.260×10 ⁻⁹	62.76	6.457×10 ⁻⁴	36.1	0.5				
Ankerite		9.8	1.260×10 ⁻⁹	62.76	6.457×10 ⁻⁴	36.1	0.5				
Dawsonite		9.8	1.260×10 ⁻⁹	62.76	6.457×10 ⁻⁴	36.1	0.5				
Ca-smectite		151.6	1.660×10 ⁻¹³	35	1.047×10 ⁻¹¹	23.6	0.34	3.020×10 ⁻¹⁷	58.9	-0.4	
Pyrite		12.9	k ₂₅ =2.818×10 ⁻⁵ E _a =56.9 n(O ₂ (aq))=0.5		k ₂₅ =3.02×10 ⁻⁸ E _a =56.9 n(H ⁺)=-0.5, n(Fe ³⁺)=0.5						

LBNL Simulator TOUGHREACT

Processes:

- Multiphase fluid and heat **flow**: TOUGH2 V2 (Pruess, et al., 1999)
- **Transport**: advection and diffusion in both liquid and gas phases
- Chemical **reactions**:
 - Aqueous complexation
 - Acid-base
 - Redox
 - Mineral dissol./precip. (equilibrium and/or kinetics)
 - Gas dissol./exsol.
 - Cation exchange
 - Surface complexation
 - Linear K_d adsorption
 - Decay

Features:

- Changes in porosity and permeability, and unsaturated zone properties due to mineral diss./ppt. and clay swelling
- Gas phase and gaseous species are active in flow, transport, and reaction
- General: Porous and fractured media; 5 ϕ -k models; rate laws; any number of chemical species
- Wide range of conditions: P, T, pH, Eh, Salinity
- <http://esd.lbl.gov/TOUGHREACT/>
- (Xu et al., 2004)



TOUGHREACT (2)

Equations for fluid and heat flow, and chemical transport.

General governing equations:
$$\frac{\partial M_{\mathbf{k}}}{\partial t} = -\nabla F_{\mathbf{k}} + q_{\mathbf{k}}$$

Water: $M_w = \phi(S_l \rho_l X_{wl} + S_g \rho_g X_{wg})$ $F_w = X_{wl} \rho_l u_l + X_{wg} \rho_g u_g$ $q_w = q_{wl} + q_{wg}$

Air: $M_a = \phi(S_l \rho_l X_{al} + S_g \rho_g X_{ag})$ $F_a = X_{al} \rho_l u_l + X_{ag} \rho_g u_g$ $q_a = q_{al} + q_{ag}$

Heat: $M_h = \phi(S_l \rho_l U_l + S_g \rho_g U_g) + (1-\phi) \rho_s U_s$ $F_h = \sum_{\beta=l,g} h_{\beta} \rho_{\beta} u_{\beta} - \lambda \nabla T$ q_h

where
$$u_{\beta} = -k \frac{k_{\phi}}{\mu_{\beta}} (\nabla P_{\beta} - \rho_{\beta} \mathbf{g}) \quad \beta = l, g \quad (\text{Darcy's Law})$$

Chemical components in the liquid phase ($j = 1, 2, \dots, N_l$):

$$M_j = \phi S_l C_{jl} \quad F_j = u_l C_{jl} - D_l \nabla C_{jl} \quad q_j = q_{jl} + q_{js} + q_{jg}$$

Chemical components in the gas phase ($k = 1, 2, \dots, N_g$):

$$M_k = \phi S_g C_{kg} \quad F_k = u_g C_{kg} - D_g \nabla C_{kg} \quad q_k = -q_{kg}$$

where $C_{kg} = f_{kg} / RT$ (gas law)

Example of chemical reaction equations

General dissociation reactions $S_i^s = \sum_{j=1}^{N_C} \nu_j S_j^p$

(1) General mass action equations: $K_i a_{S_i^s} = \sum_j (a_{S_j^p})^{\nu_j}$

Aqueous dissociation: $\text{HCO}_3^- = \text{CO}_3^{2-} + \text{H}^+$ $K_{\text{HCO}_3^-} \gamma_{\text{HCO}_3^-} c_{\text{HCO}_3^-} = \gamma_{\text{CO}_3^{2-}} c_{\text{CO}_3^{2-}} \gamma_{\text{H}^+} c_{\text{H}^+}$

Mineral dissolution: $\text{CaCO}_3(s) = \text{CO}_3^{2-} + \text{Ca}^{2+}$ $K_{\text{CaCO}_3(s)} = \gamma_{\text{Ca}^{2+}} c_{\text{Ca}^{2+}} \gamma_{\text{CO}_3^{2-}} c_{\text{CO}_3^{2-}}$

Gas dissolution: $\text{CO}_2(g) = \text{CO}_2(aq)$ $K_{\text{CO}_2(g)} f_{\text{CO}_2(g)} = \gamma_{\text{CO}_2(aq)} c_{\text{CO}_2(aq)}$

(2) Rate expressions: $R_m = \sigma_m k_m (1 - \Omega_m^*)^p$ negative for precipitation

Calcite dissolution rate (first order): $r_{\text{CaCO}_3(s)} = k_{\text{CaCO}_3(s)} A \left(1 - \frac{Q_{\text{CaCO}_3(s)}}{K_{\text{CaCO}_3(s)}} \right)$

$Q_{\text{CaCO}_3(s)} = \gamma_{\text{Ca}^{2+}} c_{\text{Ca}^{2+}} \gamma_{\text{CO}_3^{2-}} c_{\text{CO}_3^{2-}} = K_{\text{CaCO}_3(s)}$ at equilibrium

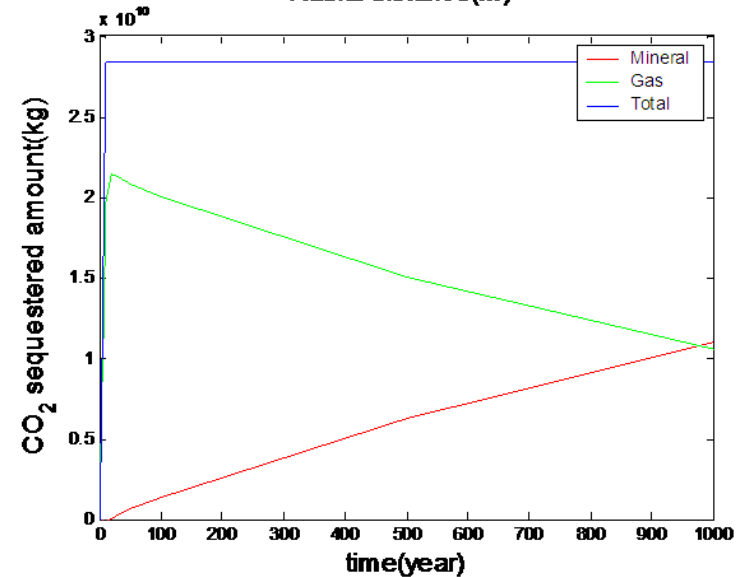
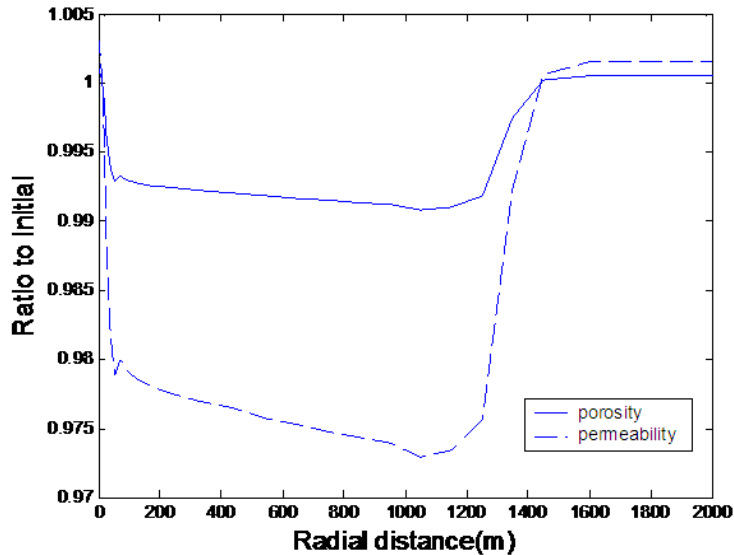
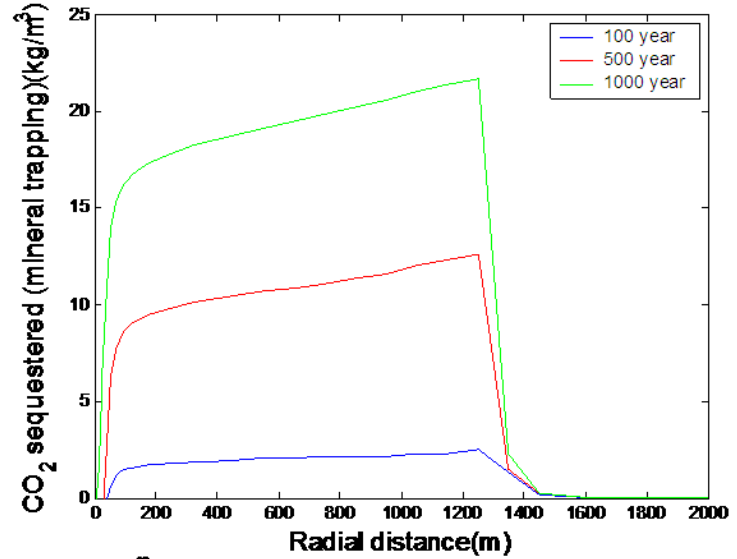
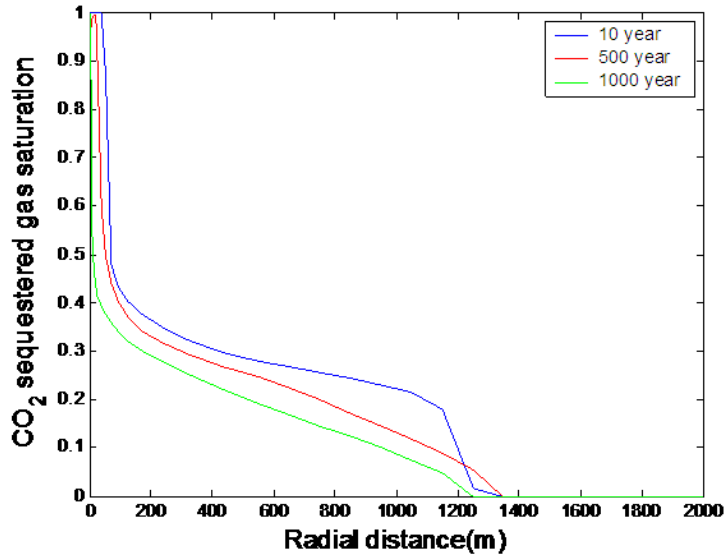
(3) Conservation of chemical component in a closed chemical system:

Carbonate component CO_3^{2-} : $T_{\text{CO}_3^{2-}} = C_{\text{CO}_3^{2-}} + c_{\text{CO}_2(g)} + c_{\text{CaCO}_3(s)}$

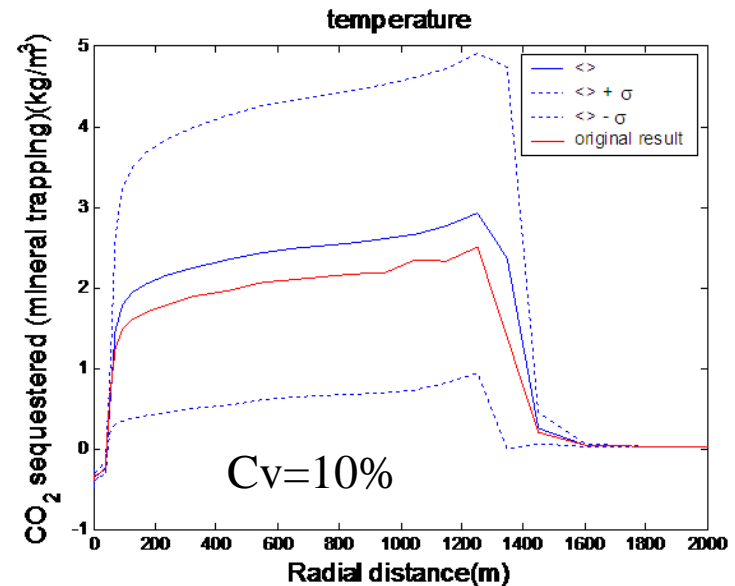
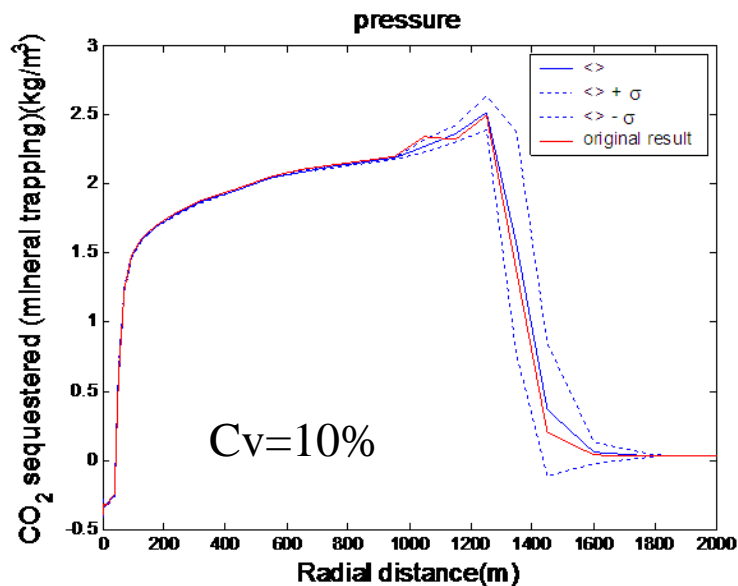
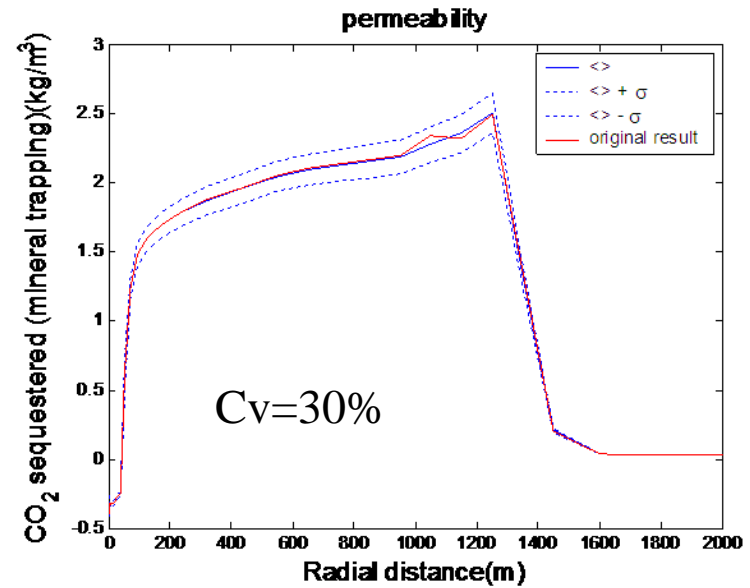
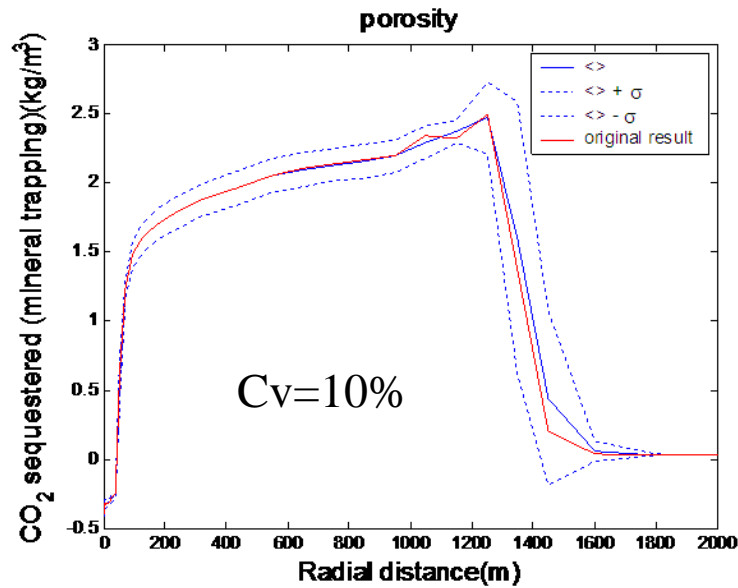
where $C_{\text{CO}_3^{2-}} = c_{\text{CO}_3^{2-}} + c_{\text{HCO}_3^-} + c_{\text{CO}_2(aq)}$ (total dissolved, subject to transport)

TOUGHREACT Simulation Results

Heli Bao, USC

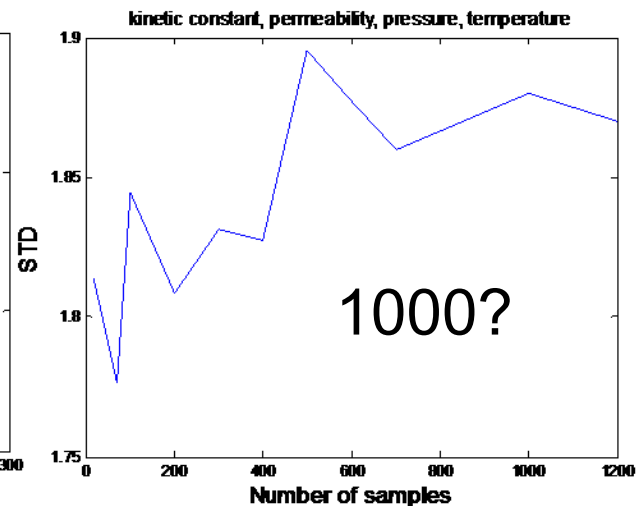
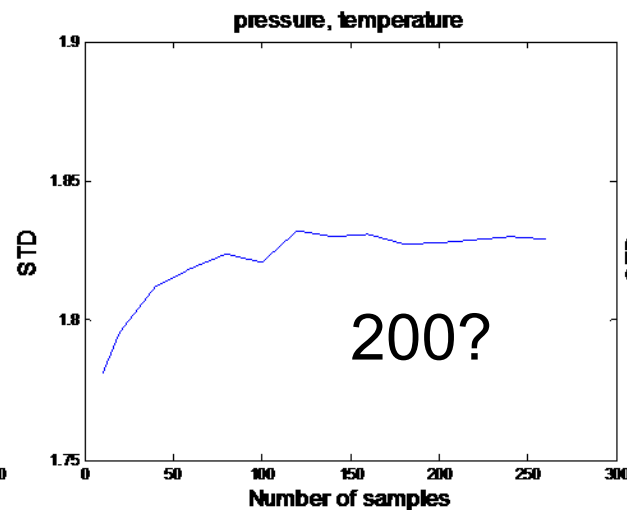
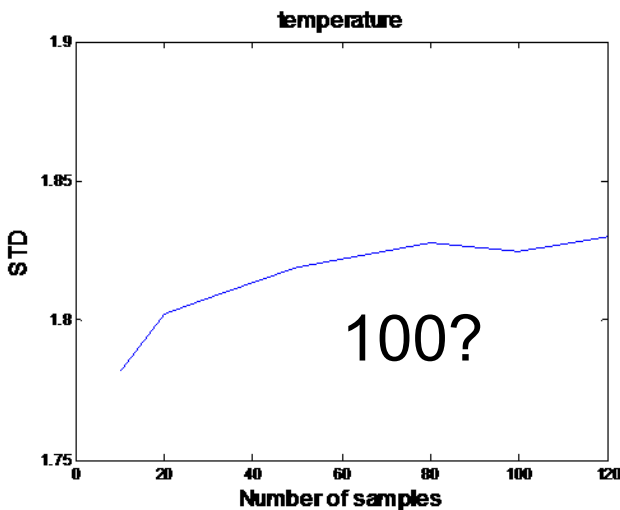


Effects of Variability: Single Parameter



Effects of Variability: Multiple Parameters (1)

The number of Latin Hypercube samples needed to obtain stable variance of sequestered CO₂:



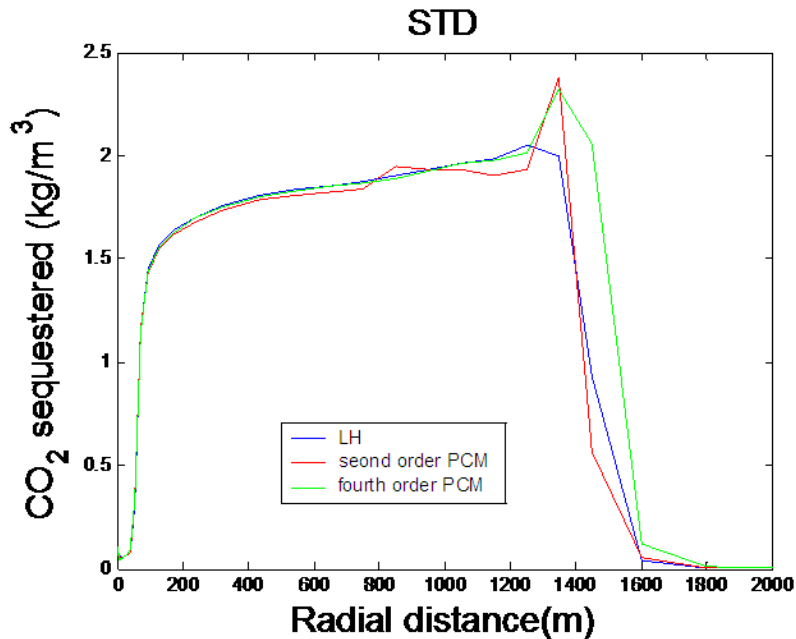
Questions:

How many samples are needed for 10s -100s parameters?

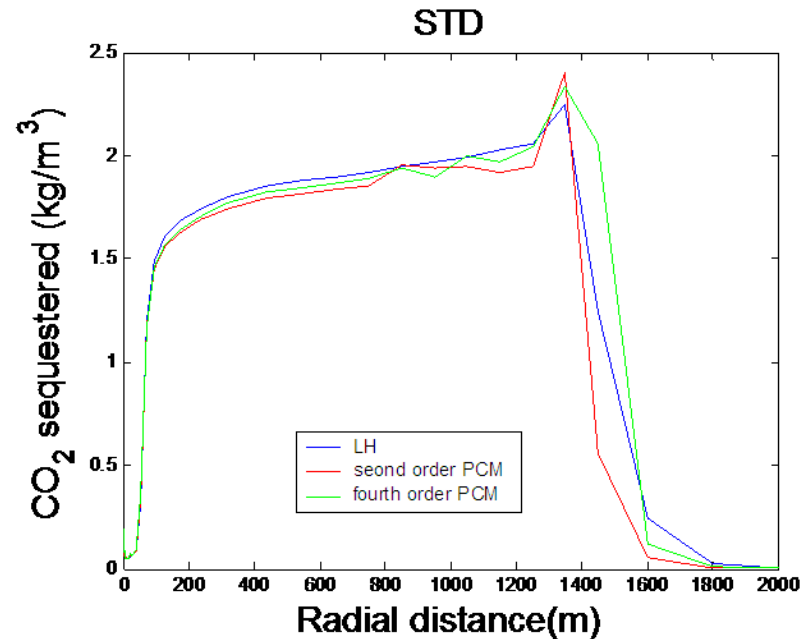
LHS or direct sampling MC?

How do the uncertainties interact among them?

Effects of Variability: Multiple Parameters (2)



P, T: $C_v=10\%$
LH: 200 runs
PCM-2nd: 6 runs
PCM-4th: 15 runs

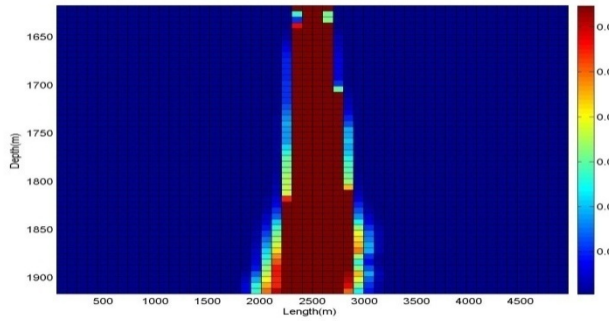


K_{Chlorite} , P, T, k: $C_v=10\%$
LH: 1000 runs
PCM-2nd: 15 runs
PCM-4th: 70 runs

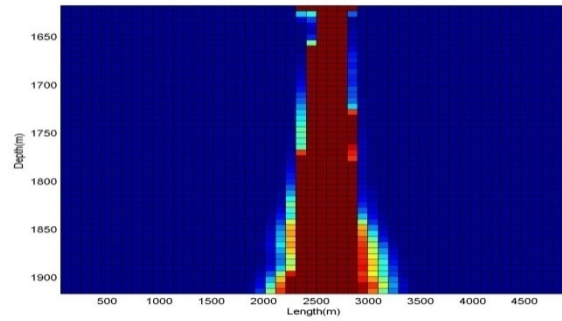
LH: Latin Hypercube sampling

PCM: Probabilistic Collocation Method (Li and Zhang, 2007)

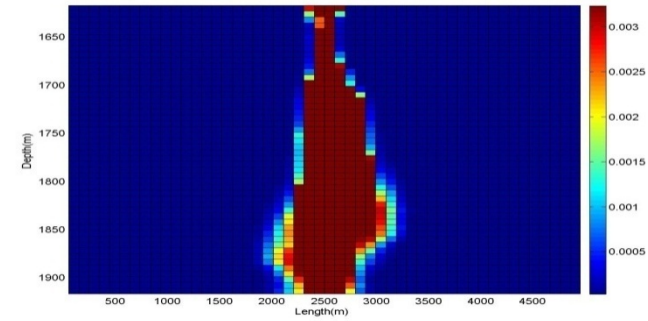
Plumes and Averages



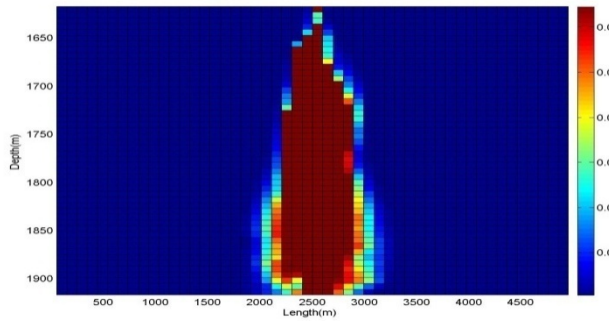
Sample 1



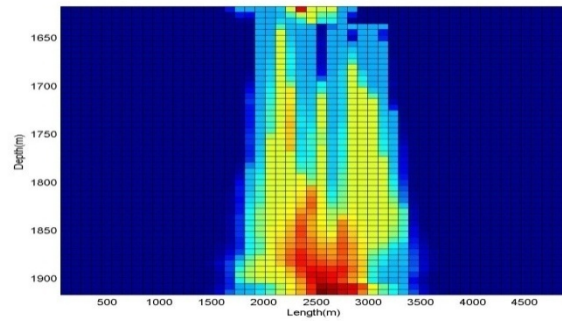
Sample 2



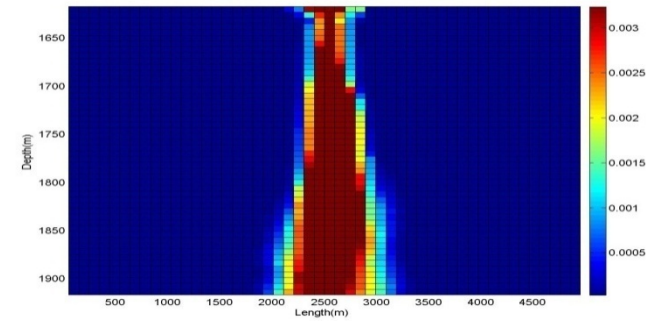
Sample 3



Sample 4

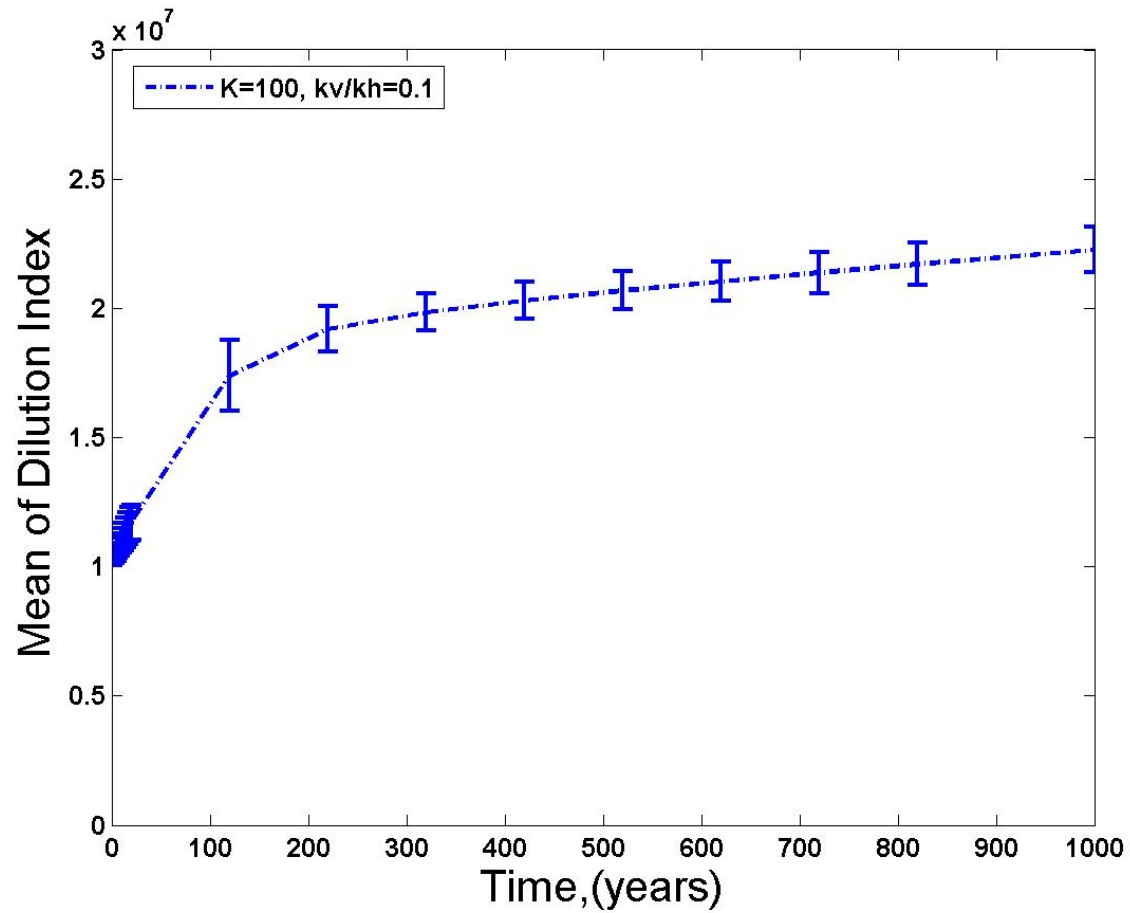


Absolute Dispersion
Averaging

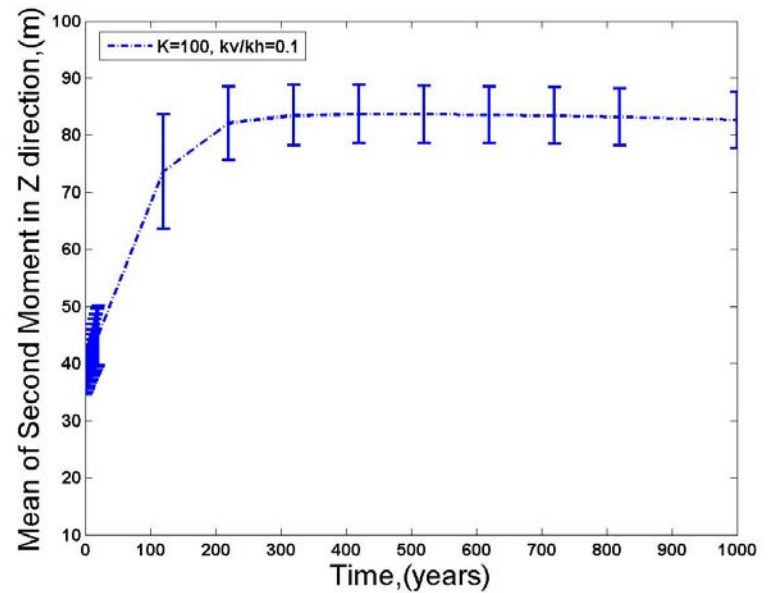
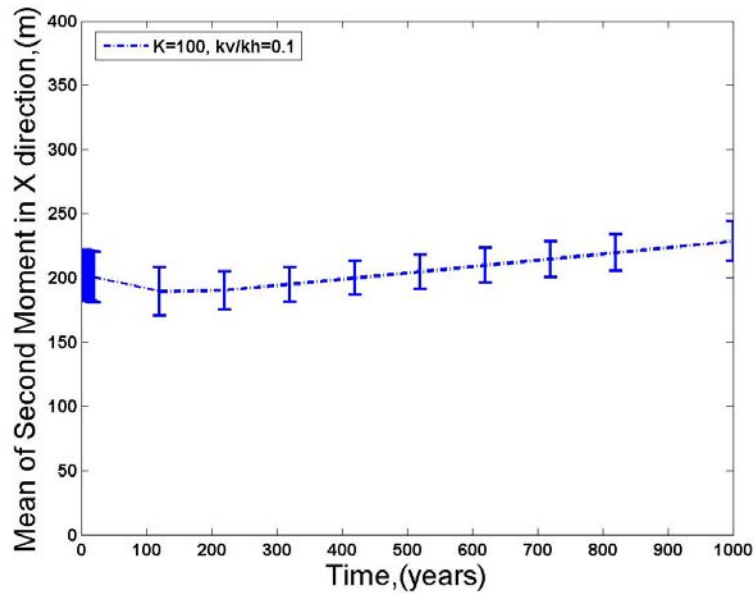
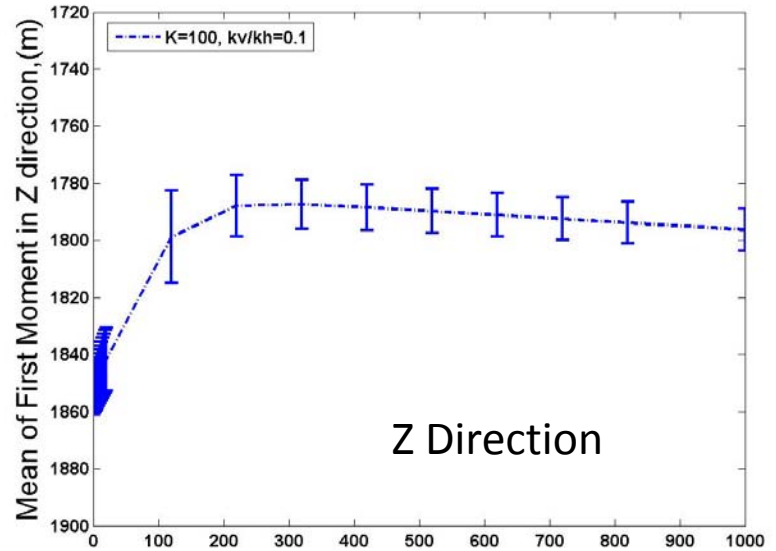
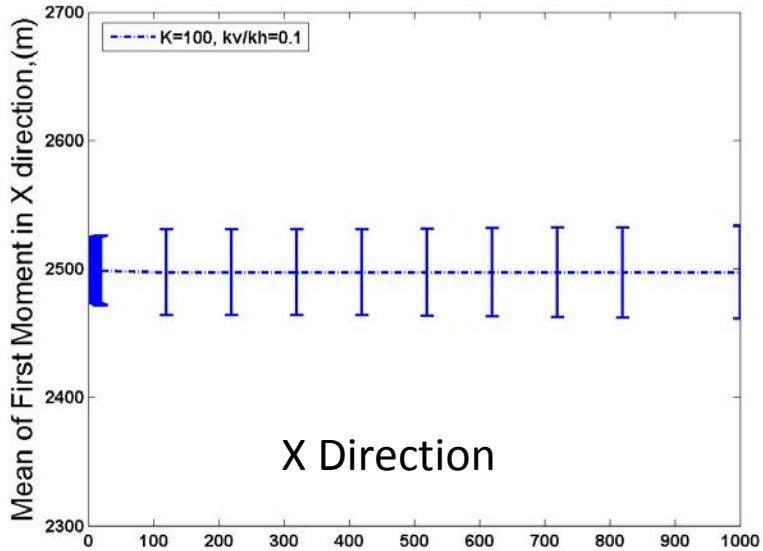


Relative Dispersion
Averaging

Dilution Index



Spatial Moments



Challenges in Uncertainty Quantification for GCS

- Multiscale and multiphysics
- Highly nonlinear
- Time consuming for each simulation
 - CPU: Several minutes to several days for each
- Multiple parameters
 - 10s to 100s
 - Various distributions
 - Some being random fields
 - Approximations with finite dimensionality (N)
 - N could be 100s or 1000s (when with multiple fields)
 - MC or PCE: $M = \frac{(N+d)!}{N! d!}$ (d - order of polynomials)?

Key Ingredients for Viable UQ Approaches

- Accurate and robust
- Efficient, in particular, for large-scale problems
- Compatible with existing simulators
 - ✓ Non-intrusive

Summary

- Geological sequestration presents an immediate, low-cost option for carbon management.
- Carbon sequestration is an important measure for sustaining fossil fuel based economy.
- In spite of past experiences, many fundamental R&D issues are outstanding:
 - Prediction under uncertainties: Development, validation, and verification
 - Monitoring and verification technologies
 - Performance and risk assessment
 - Public awareness and acceptance
- The field of carbon sequestration is still in its infancy --- providing ample research opportunities